

Alexandria, Egypt, before Alexander the Great: A multidisciplinary approach yields rich discoveries

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ABSTRACT

Historic records refer to Rhakotis as a settlement on Egypt's Mediterranean coast before Alexander the Great founded the famous Mediterranean port city of Alexandria in B.C. 332. Little is known of Rhakotis, however, because the site has yet to be clearly identified beneath the modern city. This problem motivated a geoarchaeological investigation of sediment cores from Alexandria's East Harbor, from which radiocarbon-dated sections of pre-Alexander age (>2300 yr B.P.) have been obtained for study. These core sections comprise a number of critical components, five of which are emphasized here: ceramics, rock fragments derived from Middle and Upper Egypt, and sediment with markedly increased contents of lead, heavy minerals, and organic matter. A multidisciplinary approach, by which archaeological, stratigraphical, petrological, and geochemical methodologies are applied to study the five distinct core components, reaffirms that a sum can be greater than its parts. Together, the diverse markers in the dated core sections enable us to confirm human activity to at least seven centuries before B.C. 332 on the mainland coast, where Alexandria would later be established. Alexander's city, it now appears, rose from a pre-existing town whose inhabitants had long before recognized the favorable harbor potential of this Egyptian coastal sector. The discoveries, providing direct evidence of the settlement's early (to ca. B.C. 1000) existence, are intended to prompt new exploratory efforts on land and offshore to further delineate that center's actual position and history.

INTRODUCTION

Coastal and marine geology are now well-established earth science subfields that developed from the turn of the twentieth century to the 1950s. Study of archaeological vestiges submerged in the marine realm, on the other hand, is a more recent scholarly pursuit, having evolved primarily after World War II. Integration of geological and archaeological sciences to investigate offshore sites has emerged in close association with the improvement of diving technology and equipment for offshore exploration. These include enhanced underwater

drilling, photography, and television, along with refinement of applicable high-resolution seismic methodologies and surveys by research submarine and remote operated vehicle. Coastal geoarchaeology reached a subdiscipline threshold ~25 years ago, at the time of publication of the multi-authored volume on Quaternary coastlines and marine archaeology edited by Masters and Flemming (1983). Since then, the number of studies that emphasize integration of varied geological and archaeological approaches in the marine realm has progressively risen. Of special note is the increased use of a classic geological methodology, sediment coring, to help resolve archaeological problems at sites that presently lie beneath the waves. This sub-bottom technology has been applied with successful results in most world oceans, especially in the Mediterranean (Morhange et al., 2005; Marriner and Morhange, 2007; Stanley, 2007).

The present investigation integrates archaeological, geological, and geochemical data obtained from sediment cores that provide evidence of early human activity in the Alexandria region of Egypt. The focus is on identifying new information dated to well before the arrival of Alexander the Great, who founded this major Mediterranean port city in B.C. 332 (Fig. 1).

BRIEF BACKGROUND

Historians generally agree that Rhakotis, or Râ-Kedet, was a settlement established before the fourth century B.C. in the area subsequently developed as Alexandria. Rhakotis has been vaguely alluded to as a modest fishing village of little significance, a more substantial walled center, or possibly a fortified settlement (Fraser, 1972; Empereur, 1998; Baines, 2003; McKenzie, 2003; Ashton, 2004). The modern city of Alexandria, with nearly four million inhabitants and an extensive cover of municipal and industrial construction, has almost entirely buried the remains of earlier habitation (Empereur, 1998). Although a city area south of the Heptastadion (Fig. 2) is called Rhacotis (Rowe, 1954), no archaeological excavation to date has revealed the presence of an early pre-Alexandrian site.

A record of early nautical activity near Pharos Island, positioned ~1 km seaward of the Alexandria mainland (Fig. 2), was initially passed down as oral history (from ca. B.C. 1200–1100) and then centuries later (B.C. 800–750) was recorded in Homer's epic *Odyssey*: "Now in the surging sea an island lies, – Pharos they call it, – By it there lies a bay with a good anchorage, from which they send the trim ships off to sea and get them drinking water." After Homer, historians and geographers intermittently refer to this Egyptian sector in their texts, such as Herodotus' *The History* (fifth century B.C.) and Strabo's *Geographia* (near the turn of the first century A.D.). Subsequently, scholars surmise that Minoan, Philistine, Phoenician, ancient Greek, and other early mariners sailing the eastern Mediterranean sought protection in the lee of Pharos Island long before Ptolemaic (B.C.

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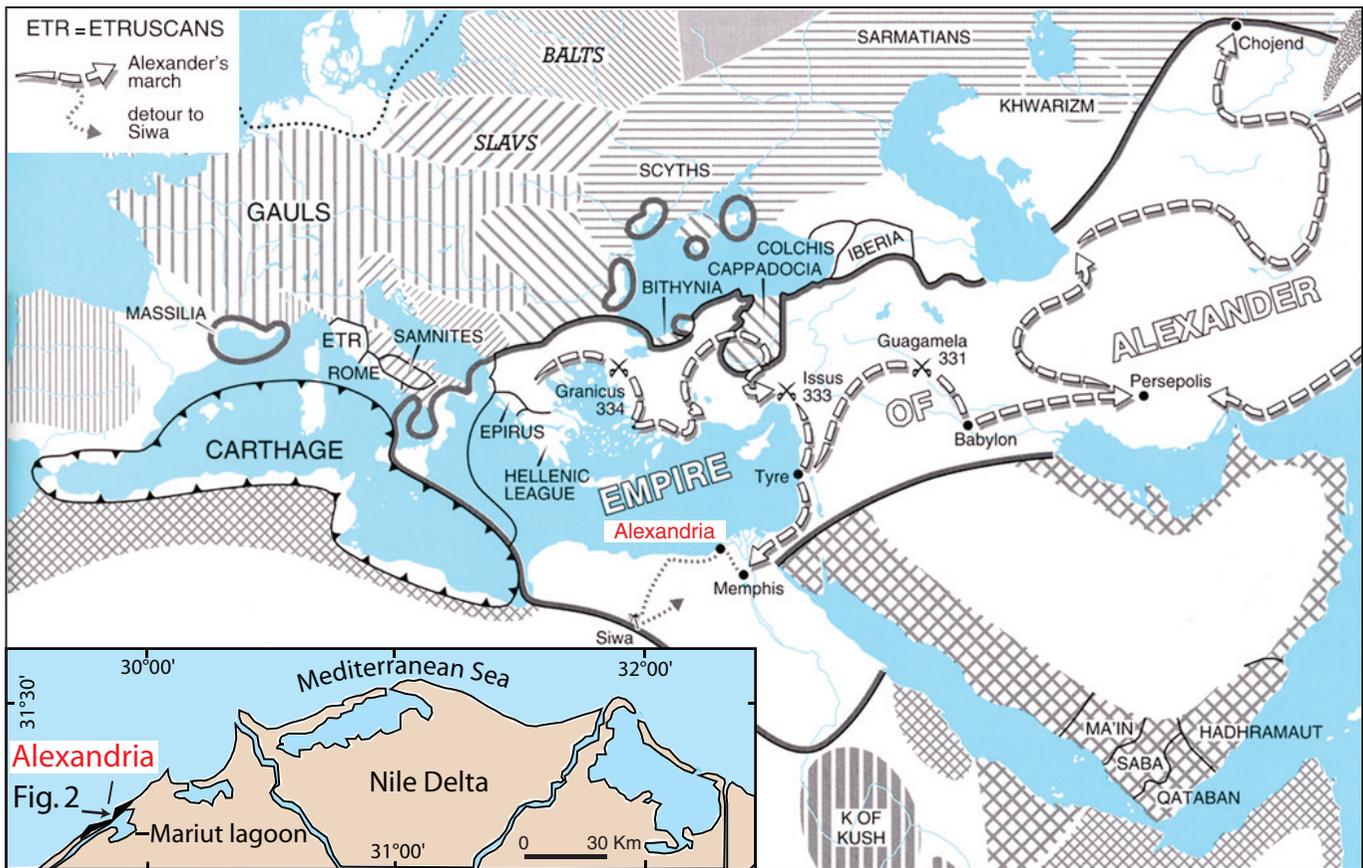


Figure 1. Map showing the geographic extent of Alexander the Great's empire and the Ptolemaic reign (post-B.C. 332) in the eastern Mediterranean (modified after McEvedy, 2002). Inset shows the Alexandria, Egypt, study area.

332–30) rule (Jondet, 1916; Weill, 1919). Additional evidence of pre-Alexander occupation in this region is provided by early archaeological sites in proximal sectors (Egypt Exploration Society, 2005), including those along the Mariut lagoon to the south (Rowe, 1954) and the Nile delta coast to the east of Alexandria (Stanley, 2005, 2007).

Recent preliminary findings record more direct evidence of settlement prior to the Ptolemies in, or proximal to, the modern city. These are based on assessments of cores collected on land adjacent to the East Harbor of Alexandria (Goiran et al., 2000; Véron et al., 2006) and in the East Harbor proper (Stanley and Landau, 2005). New findings summarized herein are obtained by a comprehensive and multidisciplinary study of dated cores recovered in the harbor.

METHODS

The present East Harbor basin covers an area of ~2.5 km² and is bound to the south by an arcuate coastline bordered by the city of Alexandria (Fig. 2). The region has been the focus of numerous investigations, including geography (Goddio et al., 1998), oceanography (Inman and Jenkins, 1984), geology and paleogeography (Warne and Stanley, 1993; Goiran et al., 2005), stratigraphy (Goiran et al., 2000), sedimentology (El-Wakeel and El-Sayed, 1978; Wali et al., 1994), and geochemistry (Véron et al., 2006).

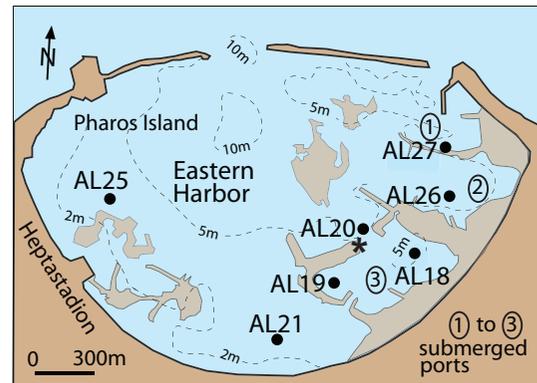


Figure 2. Configuration of the Eastern Harbor and locations of seven vibracore sites off Alexandria. Light brown features in the Eastern Harbor are shallow reefs and submerged port structures (after Goddio et al., 1998). Asterisk denotes eastern end of now-subsided Antirrhodos Island, where late fifth and early fourth century B.C. wood was recovered.

New information to define this basin's early Holocene to present evolution is provided by seven vibracores (lengths ~2.0–5.5 m; Fig. 3) collected in the East Harbor (Stanley and Bernasconi, 2006). The 52 Accelerator Mass Spectrometry (AMS) dates obtained for core samples are given in uncalibrated ¹⁴C

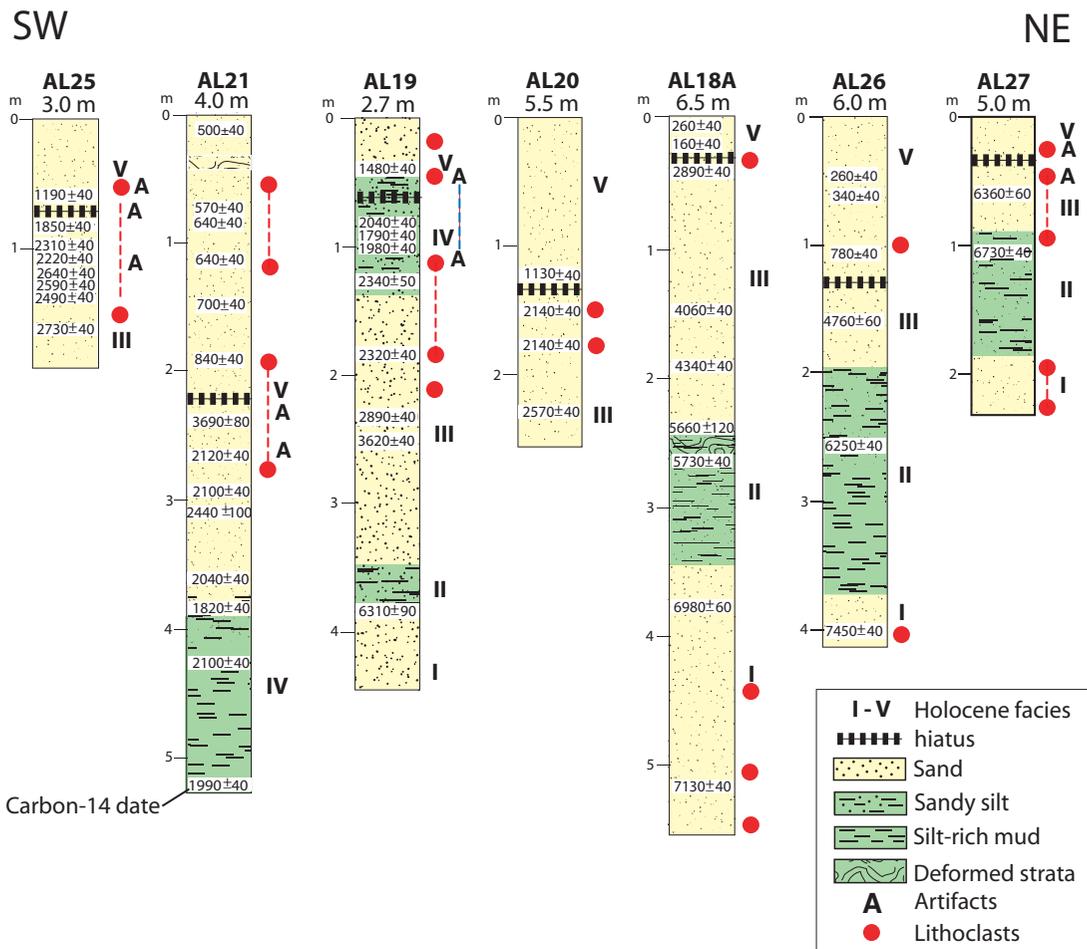


Figure 3. Stratigraphic logs of the seven vibracores collected in the East Harbor, showing five dominant lithofacies (I–V), conventional radiocarbon (uncalibrated) dates (in thousand yr B.P.), and positions of artifacts and lithoclasts. Modified after Stanley and Bernasconi (2006).

radiocarbon years (most shown in Fig. 3). Samples ($n = 441$) were taken throughout the sand-rich cores at <15 cm intervals for petrologic study. Lithological logs of the cores and results of different analyses are given in Stanley and Bernasconi (2006).

Lead isotopic analyses were performed at the Carnegie Institution of Washington. After addition of a ^{205}Pb spike, ~ 20 – 50 mg of coarse core material was dissolved, and the Pb was separated and analyzed using procedures described in Carlson et al. (2006). Some samples, labeled with an “L” in Table 1, were first subjected to an acetic acid leach. In most cases, $>50\%$ of the sample was dissolved in the acetic acid. After leaching, the residue was dissolved as described above, while the leach was first dried and then redissolved in HBr in preparation for Pb separation. All measured Pb isotopic compositions are corrected for mass fractionation based on these average standard values compared to the isotopic composition of NBS 981 reported by Todt et al. (1996).

OBSERVATIONS

Core Content in Greek to Recent Time

Stratigraphic analysis of dated borings identifies five mostly carbonate lithofacies (coded I to V) from base to top of East

Harbor cores (Figs. 3 and 4A): Lower Sand (I) and Lower Mud (II) units of early to mid-Holocene age (>7500 to ca. 5600 yr B.P.); Middle Sand (III) facies of mid- to late Holocene age (ca. 5600 – 2300 yr B.P.); and Upper Mud (IV) and Upper Sand (V) of late Holocene age since Ptolemaic rule (<2300 yr B.P.). Core AL19 comprises all five stratigraphic facies, with a basal section older than 6310 ± 90 radiocarbon yr B.P., and thus serves as a representative, or type, boring to define changes through time denoted by diagnostic archaeological, geological, and geochemical criteria.

Upper Sand (V) and Upper Mud (IV) sections in this and other East Harbor cores comprise the most abundant and diverse suites of archaeological material. These upper two units contain potsherds, pebble-size rock clasts, and high concentrations of heavy minerals (Fig. 4D), lead (Fig. 5A), organic matter, quartz, and crystalline and aggregate limestone, when compared to the three older facies (I–III). Rapid municipal development during the reign of the Ptolemies and Romans gave rise to marked anthropogenic signals in the two upper lithofacies. For example, very high lead concentrations (>100 ppm) occur in post-Alexander sediment in core AL19 (Fig. 5A); this is also recorded in core sections of this age collected on land in Alexandria proper (Goiran et al., 2000; Véron et al., 2006). The

findings denote heavy lead use by Greeks and Romans during Alexandria's swift expansion.

Accelerated construction along the mainland shore and in the East Harbor proper (Fig. 2) during this time (IV–V) is amply recorded (Empereur, 1998; Goddio et al., 1998; Hesse, 1998), including the building of the Heptastadion, the large freshwater aqueduct-causeway system between the city and Pharos Island (Hesse, 1998; Goiran et al., 2000, 2005). Increased organic matter content in upper core facies IV and V includes fibers of *Phragmites* sp. and algae. Such materials were derived from the city's sewage runoff and from Alexandria Canal discharge into the harbor (Stanley and Bernasconi, 2006).

Key Markers in Pre-Greek Time

Until now, the record of human activity in Alexandria prior to the Ptolemies has been sorely limited. Underwater diver excavation at the eastern end of now-submerged Antirrhodos Island (marked by an asterisk in Fig. 2) recovered posts of elm (*Ulmus* sp., a non-local wood) radiocarbon-dated (calibrated calendar years) at B.C. 410 ± 40 and planks of pine (*Pinus* sp.) at B.C. 395 ± 40 (Goddio et al., 1998).

Materials found in cores recovered in the East Harbor, however, provide more ample evidence of older pre-fourth century B.C. (>2400 yr B.P.) human activity, especially in sediment forming the upper part of the Middle Sand (III) unit in East Harbor cores (Fig. 4A). Foremost are potsherds of early age that were recovered in this part of unit III in core AL25, in the western East Harbor (Fig. 3). Ceramic fragments include coarse and poorly fired material, mostly cooking vessels, although some thinner ceramics from bowls and small jars are also observed. The pottery is wheel-made and appears to be of local production rather than imported. Archaeological analysis of ceramic fragments, including ones preserved with slip (Fig. 4B, 1–2) and others that are rimmed (Fig. 4B, 3–4), show they are most closely comparable to typical southeastern Mediterranean ware made in the ninth to seventh centuries B.C. (plates 81–83 in Tufnell, 1953; National Museum of Natural History collections). These potsherds and lithic fragments prevail in sediment radiocarbon dated to well before 2330 yr B.P. (calibrated dates of B.C. 940 to B.C. 420).

Most cores include carbonate pebbles derived from exposures proximal to the East Harbor that originated in the lower (older) sections of cores AL18A, AL26, and AL27. Core AL19 also contains pebble-sized rock fragments (diorite, gabbro, quartzite, marble, and dense fossiliferous limestone) in the upper half of facies III at a depth of 2.3 m from core top (Fig. 4C). These lithoclasts, with diameters of 1–5 cm, are of non-local (allochthonous) derivation and were obtained from distant quarries, mostly in Middle and Upper Egypt. This upper facies III section, positioned >50 cm beneath the base of the Upper Mud (IV) unit, is radiocarbon dated as older than ca. 2320 yr B.P. but younger than 2890 ± 40 yr B.P. (Fig. 4A). Rock reached Alexandria's East Harbor either by vessel or overland transport for use in building structures; in the case of core site AL19, located ~500 m from the present coastline (Fig. 2), this material was likely brought to localities near the southern East Harbor margin. Some clasts may also have been derived from materials recycled from statuary or other artifacts of Egyptian

TABLE 1. Pb CONCENTRATION AND ISOTOPIC COMPOSITION OF ALEXANDRIA CORE SEDIMENTS AND ARTIFACTS

Sample no.	Depth (cm)	Sample	[Pb] ppm	204/206	207/206	208/206
TJ1	56	Clayey silt	171	0.05401	0.84601	2.08725
TJ2	96	Clayey silt	286	0.05409	0.84785	2.09116
TJ9	120	Shelly silt	102	0.05367	0.84034	2.07878
TJ24	123	Carbonate sand	84	0.05395	0.84481	2.08589
TJ25	126	Carbonate sand	78	0.05362	0.84187	2.08423
TJ16	129.5	Silty carbonate sand	49	0.05363	0.84263	2.08632
TJ17	139	Silty carbonate sand	48	0.05370	0.84271	2.08528
TJ18	149.5	Silty carbonate sand	65	0.05350	0.84089	2.08428
TJ3	159	Dark clay	68	0.05341	0.83923	2.08245
TJ10	183	Silt w/shell frag.	83	0.05330	0.83602	2.07573
TJ10L	183	Silt w/shell frag.	74	0.05330	0.83621	2.07632
TJ11	204.5	Fine silt	56	0.05318	0.83394	2.06831
TJ11L	204.5	Fine silt	51	0.05319	0.83388	2.06732
TJ12	226	Silt w/shell & rock frags.	52	0.05314	0.83391	2.06727
TJ12L	226	Silt w/shell & rock frags.	43	0.05315	0.83338	2.06554
TJ26	230	Silty sand	16.7	0.05318	0.83566	2.07105
TJ27	235	Silty carbonate sand	15.9	0.05339	0.83691	2.07127
TJ28	240	Silty carbonate sand	13.6	0.05335	0.83762	2.07550
TJ13	243	Silt w/shell frags.	37	0.05367	0.84127	2.08179
TJ13L	243	Silt w/shell frags.	22	0.05387	0.84342	2.08225
TJ19	245	Clayey sand	8.2	0.05376	0.84249	2.07925
TJ20	247	Clayey sand	12.1	0.05374	0.84323	2.08363
TJ21	249	Clayey sand	10.4	0.05365	0.84089	2.07847
TJ4	251	Clayey silt	1.99	0.05397	0.84348	2.07967
TJ29	273	Silty carbonate sand	0.66	0.05321	0.83064	2.05908
TJ30	326	Sandy silt carbonate	0.74	0.05312	0.83082	2.05704
TJ5	349	Silty sand	4.36	0.05323	0.83250	2.06064
TJ31	405	Silty carbonate sand	1.12	0.05314	0.82924	2.05593
<u>Archaeological samples</u>						
TJ6		Metal artifact, Core AL25	6.65	0.05425	0.84835	2.08671
TJ14		Ceramic glaze, Core AL25	49	0.05391	0.84389	2.08343
TJ15		Ceramic no glaze, Core AL25	17.9	0.05341	0.83703	2.07152
TJ22		Ceramic, Core AL25	17.1	0.05349	0.83855	2.07631
TJ23		Ceramic, Core AL25	128	0.05356	0.84093	2.08351
TJ32		Pb crab	30.4*	0.05325	0.83646	2.07953
TJ33		Wooden coffin lid	83	0.05341	0.83182	2.05998
TJ34		Clay cartouche	206	0.05389	0.84465	2.08873
TJ35		Terra cotta figurine	18.6	0.05366	0.83853	2.07095
TJ36		Mummy case	190	0.05767	0.87327	2.10021
TJ37		Clay vase	17.0*	0.05029	0.79379	2.00047

Note: Artifacts TJ32–TJ37 from the Smithsonian Institution's Egyptian and Nile delta materials collection. Pb concentrations determined by isotopic dilution using a ²⁰⁵Pb spike. Uncertainties on Pb concentration: <1%. Isotopic compositions have uncertainties of ^{204/206}Pb = 0.06%, ^{207/206}Pb = 0.10%, and ^{208/206}Pb = 0.12%. Sample numbers ending in "L" are acetic acid leaches; residues of leaching are denoted by the same sample number without the "L" ending.

*wt%.

dynastic origin. Polished, flat, well- to very well-rounded clasts are typical of rock worn on a beach or in a shallow aqueous setting subject to strong swash action (Stanley, 2005).

Similarly, in the mid- to upper section of the Middle Sand (III) unit there is a marked increase in proportions of heavy minerals (to 1%; Fig. 4D) and organic matter content (to 1.5% by weight; Stanley and Landau, 2007). As with potsherds and rock clasts, the increased content of heavy minerals and organic matter found in carbonate-rich facies dated prior to Ptolemaic rule are associated with increased human activity. Construction activity involved use of noncarbonate rock material and sediment transported to the area, while organic matter

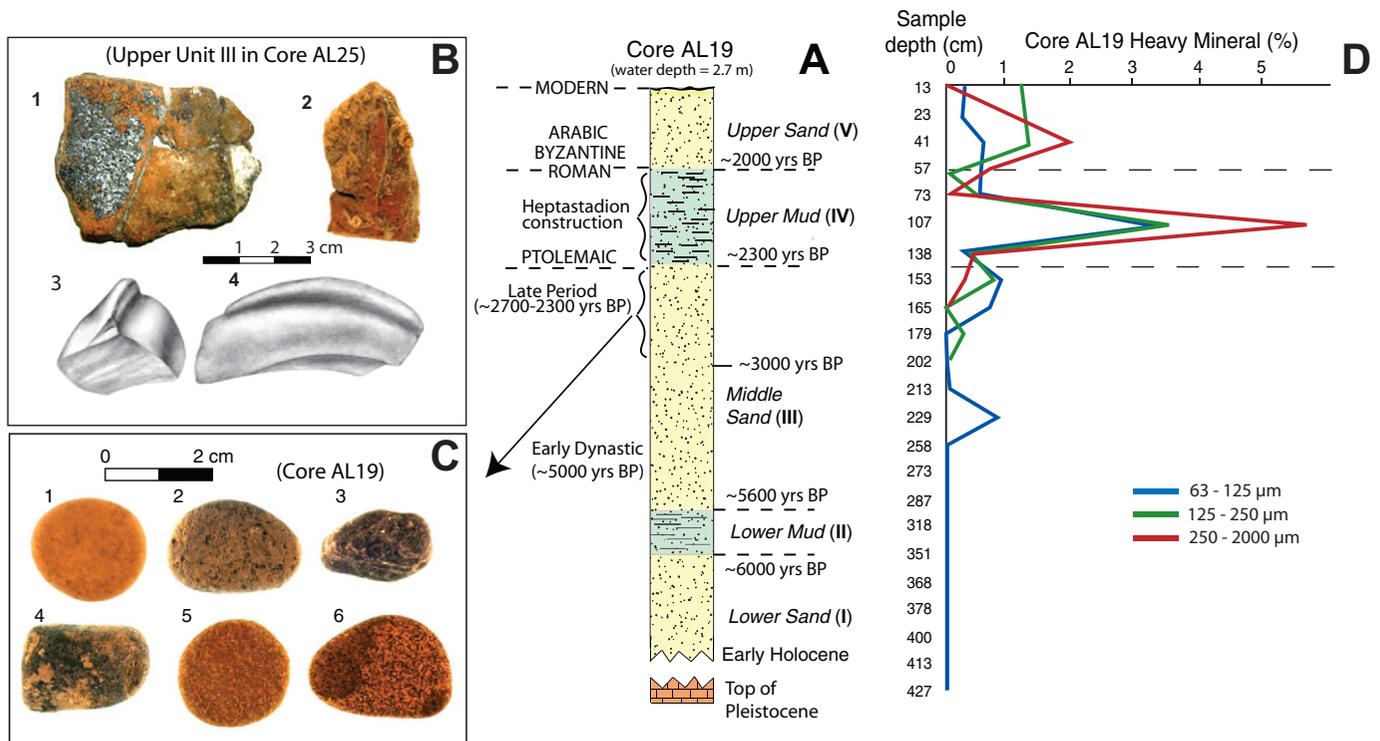


Figure 4. (A) Lithologic log of type core AL19 showing five (I–V) dated stratigraphic units. (B) Sherds of representative ware locally produced and dated from ninth and eighth centuries B.C.: slip (1, 2) and rimmed (3, 4). (C) Examples of pebbles of diverse lithologies derived from Middle and Upper Egypt, dated from ca. 3000–2300 yr B.P., recovered in upper part of unit III in core AL19. 1, 2—fossiliferous limestone; 3, 4—quartzite; 5—diorite; 6—gabbro. (D) Relative percentage of heavy mineral content in very fine, fine, and medium sand-size fractions, with marked increase in upper part of unit III before 2300 yr B.P.; major peak occurs in unit IV during Alexandria’s development by Greeks and Romans.

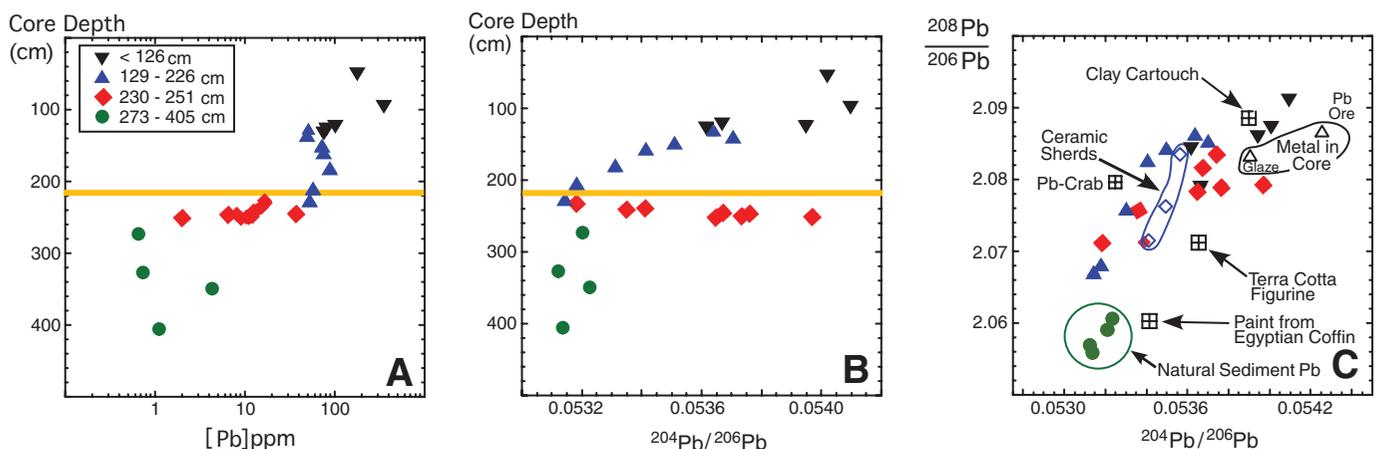


Figure 5. (A) Lead concentration, in ppm on a log-scale. (B) Isotopic ratio ($^{204}\text{Pb}/^{206}\text{Pb}$) versus depth in AL19; horizontal yellow lines denote ca. 2300 yr B.P. Alexandria timeline. Major changes in both concentration and isotopic ratio occur between ca. 2900 and 2300 yr B.P. Samples color-coded by depth in core. (C) Isotopic ratios ($^{208}\text{Pb}/^{206}\text{Pb}$ versus $^{204}\text{Pb}/^{206}\text{Pb}$) of sediment in core AL19 and artifactual material, including ceramic and metal from core AL25 and pre-Alexandria artifacts in the Smithsonian Institution’s Lower Egypt collections (see Table 1). Ceramic fragments in the core are identified by open blue diamonds. Artifactual materials are shown by cross-hatched boxes. Potential representatives of anthropogenically concentrated Pb are provided by samples from cores that include a metal fragment and glaze from ceramic potsherds in core AL25 (open triangles).

was discharged into the East Harbor by increasing volumes of municipal waste water.

Geochemical Markers in Pre-Greek Time

In addition to the four archaeological and petrologic parameters, lead concentration and isotopic composition in core AL19 (see Table 1) present a formidable additional line of evidence of human activity prior to B.C. 332. Lead concentrations are consistently below 10 ppm at depths >251 cm from core top (e.g., sections in the lower part of unit III that are older than ca. 4000 yr B.P.; Fig. 5A). From 251–230 cm (deposition in pre-Alexander time), Pb concentration gradually increases, reaching 50–80 ppm through depths of 123 cm (Ptolemaic time). Above 120 cm, Pb concentrations increase to >100 ppm, marking substantial Pb input into the East Harbor following accelerated growth of Alexandria by Greeks and Romans, as also documented by Véron et al. (2006).

In all samples below 251 cm, both $^{204}\text{Pb}/^{206}\text{Pb}$ and $^{208}\text{Pb}/^{206}\text{Pb}$ ratios are slightly lower than values typical of average continental crust (e.g., $^{204}\text{Pb}/^{206}\text{Pb} = 0.0535$ and $^{208}\text{Pb}/^{206}\text{Pb} = 2.0658$; Stacey and Kramers, 1975). These values are indicative of the isotopic composition of natural Pb deposited in sediment prior to human influence. Throughout the post-Alexander portion of the core (above 226 cm depth), Pb concentrations are very high (>43 ppm) and show continually increasing $^{204}\text{Pb}/^{206}\text{Pb}$ and $^{208}\text{Pb}/^{206}\text{Pb}$ ratios upsection in the core (Fig. 5B), consistent with increasing input of pollutant Pb from ore sources (Véron et al., 2006). Potential examples of this material include the metal fragment and ceramic glaze recovered from core AL25 (Fig. 5C).

Of special note, however, is that Pb concentration and isotopic composition do not correlate at the onset of anthropogenic Pb contamination in the interval between 251 cm and 183 cm from core top. Within the 251–235 cm interval, $^{204}\text{Pb}/^{206}\text{Pb}$ ratios increase to values similar to those seen in the shallowest portions of the core, yet both Pb concentrations and $^{208}\text{Pb}/^{206}\text{Pb}$ (Fig. 5C) remain below those characterizing the shallow portions of the core. Between ~230 and 183 cm from core top, Pb concentrations increase dramatically, yet $^{204}\text{Pb}/^{206}\text{Pb}$ ratios drop to values similar to those of the natural sediment Pb, but with significantly higher $^{208}\text{Pb}/^{206}\text{Pb}$. This may indicate an initial burst of anthropogenic Pb contamination into the East Harbor involving a source of Pb different from the ore-leads characteristic of the post-Alexander era.

The core component between 251 and 183 cm depth, characterized by high $^{208}\text{Pb}/^{206}\text{Pb}$ (~2.07 to ~2.084) and moderate $^{204}\text{Pb}/^{206}\text{Pb}$ (~0.0532 to ~0.0538) is similar in isotopic composition to ceramic potsherds from core AL25 that date to pre-Alexandrian times (Figs. 4B and 5C). These ceramic fragments have only moderate Pb concentrations (17–128 ppm), at least compared to weight percent concentrations in the analyzed pigments and metals (Table 1), and thus explain the relatively modest Pb concentration increase but marked Pb isotopic variation between 251 and 230 cm depth in the core. These data suggest that the first significant anthropogenic Pb contribution to East Harbor sediments did not derive from Greek and Roman metalworking, but instead reflect deposition of clay used perhaps for both building construction and manufacturing of ceramic vessels. The area surrounding the East Harbor

is dominantly carbonate (Stanley and Bernasconi, 2006), so the clay signature is not likely to be locally (in situ) sourced. Only well into the post-Alexandria period did Pb pollution in the East Harbor shift to isotopic compositions expected for anthropogenically concentrated Pb; for example, as associated with metallurgy, paints, and pigments.

CONCLUSIONS

The more approaches utilized in an archaeological investigation, the greater the possibility of attaining robust new findings. Results from the application of diverse geological methodologies in this study provide consistent data on five distinct and diverse components in radiocarbon-dated East Harbor cores: ceramics, allochthonous rock fragments, lead concentration, heavy minerals, and organic matter, which all increase substantially in the same upper part of the Middle Sand (III) unit dated to ca. 3000 yr ago. Stratigraphically, this depositional phase clearly corresponds to one that began well before the arrival of Alexander in B.C. 332.

Together, the five archaeological, petrological, and geochemical markers provide compelling evidence of human activity dating to as much as seven centuries before the development of Alexandria by the Ptolemies. In particular, the ceramic sherds, lead isotopes, and associated data collected from harbor sediment cores indicate that a coastal population flourished in this area during Egypt's Intermediate (ca. B.C. 1000) and Late Dynastic (pre-Ptolemaic) periods.

In summary, evidence from East Harbor cores shows that Alexandria did not grow from a barren desert, but was built atop an active town that had for centuries exploited the safe harbor setting along this Egyptian coast. Beyond providing a preliminary insight into the early settlement's history, it is expected that the investigation findings will provide impetus for further geoarchaeological exploratory efforts in this historically rich region.

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REFERENCES CITED

- Ashton, S.-A., 2004, Ptolemaic Alexandria and the Egyptian tradition, in Hirst, A. and Silk, M., eds., *Alexandria, real and imagined*: London, Ashgate, p. 15–40.
- Baines, J., 2003, Appendix: Possible implications of the Egyptian word for Alexandria: *Journal of Roman Archaeology*, v. 16, p. 61–63.
- Carlson, R.W., Czmannske, G., Fedorenko, V., and Ilupin, I., 2006, A comparison of Siberian meimechites and kimberlites: Implications for the source of high-Mg alkalic magmas and flood basalts: *Geochemistry, Geophysics, Geosystems*, v. 7, no. 11, doi: 10.1029/2006GC001342.
- Egypt Exploration Society, 2005, EES Delta Survey (reference collection of archaeological sites in the Egyptian Nile Delta): <http://www.ees.ac.uk/deltasurvey/dsintro.html> (accessed 25 May 2007).
- El-Wakeel, S.K., and El-Sayed, M.Kh., 1978, The texture, mineralogy and chemistry of bottom sediments and beach sands from the Alexandria region, Egypt: *Marine Geology*, v. 27, p. 137–160, doi: 10.1016/0025-3227(78)90077-4.
- Empereur, J.-Y., 1998, *Alexandrie Redécouverte*: Paris, Stock, 253 p.
- Fraser, P.M., 1972, *Ptolemaic Alexandria* (3 vols.): London, Oxford University Press, 2136 p.
- Goddio, F., Bernand, A., Bernand, E., Darwish, I., Kiss, Z., and Yoyotte, J., 1998, *Alexandria, the Submerged Royal Quarters*: London, Periplus, 274 p.

- Goiran, J.-P., Morhange, C., Bourcier, M., Carbonel, P., and Morigi, C., 2000, Evolution des rivages d'Alexandrie à l'Holocène récent, marge occidentale du delta du Nil, Egypte: Méditerranée, v. 94, no. 1.2, p. 83–90.
- Goiran, J.-P., Marriner, N., Morhange, C., Abd El-Maguib, M., Espic, K., Bourcier, M., and Carbonel, P., 2005, Evolution géomorphologique de la façade maritime d'Alexandrie (Egypte) au cours des six derniers millénaires: Méditerranée, v. 104, no. 1.2, p. 61–64.
- Herodotus, The History, Translated by D. Greene, 1987: Chicago, The University of Chicago Press, 699 p.
- Hesse, A., 1998, Arguments pour une nouvelle hypothèse de localisation de l'Heptastade d'Alexandrie: Alexandrina, v. 1, p. 1–33.
- Homer, The Odyssey, Translated by G.H. Palmer, 2003: New York, Barnes & Noble Classics, 339 p.
- Inman, D.L., and Jenkins, S.A., 1984, The Nile littoral cell and man's impact on the coastal zone of the southeastern Mediterranean: Scripps Institution of Oceanography: Reference Series, v. 31, p. 1–43.
- Jondet, M.G., 1916, Les Ports Submergés de l'Ancienne île de Pharos: Cairo, L'Institut Egyptien, 101 p.
- Marriner, N., and Morhange, C., 2007, Geoscience of ancient Mediterranean harbours: Earth-Science Reviews, v. 80, p. 137–194, doi: 10.1016/j.earscirev.2006.10.003.
- Masters, P.M., and Flemming, N.C., eds., 1983, Quaternary Coastlines and Marine Archaeology: London, Academic Press, 641 p.
- McEvedy, C., 2002, The New Penguin Atlas of Ancient History: London, Penguin Books, 128 p.
- McKenzie, J., 2003, Glimpsing Alexandria from archaeological evidence: Journal of Roman Archaeology, v. 16, p. 35–63.
- Morhange, C., Goiran, J.-P., and Marriner, N., eds., 2005, Coastal geoarchaeology of the Mediterranean: Méditerranée, v. 104, no. 1.2, p. 1–131.
- Rowe, A., 1954, A contribution to the archaeology of the Western Desert: Bulletin of the John Rylands Library, v. 36, p. 128–145.
- Stacey, J.S., and Kramers, J.D., 1975, Approximation of terrestrial lead isotope evolution by a two-stage model: Earth and Planetary Science Letters, v. 26, p. 207–221, doi: 10.1016/0012-821X(75)90088-6.
- Stanley, J.-D., 2005, Submergence and burial of ancient coastal sites on the subsiding Nile delta margin, Egypt: Méditerranée, v. 104, p. 65–73.
- Stanley, J.-D., ed., 2007, Geoarchaeology, Underwater Excavations in the Canopic Region of Egypt: Oxford, Oxford Centre for Maritime Archaeology, 128 p.
- Stanley, J.-D., and Bernasconi, M.P., 2006, Holocene depositional patterns and evolution in Alexandria's Eastern Harbor, Egypt: Journal of Coastal Research, v. 22, p. 283–297, doi: 10.2112/04-0348.1.
- Stanley, J.-D., and Landau, E.A., 2005, Early human activity (pre-332 BC) in Alexandria, Egypt: New findings in Eastern Harbor cores: Eos (Transactions, American Geophysical Union) Fall Meeting Supplement, v. 86, no. 52, p. F1252.
- Stanley, J.-D., and Landau, E.A., 2007, New petrological findings in Eastern Harbour cores, Alexandria, in Cole, J., ed., City and Harbour, the archaeology of ancient Alexandria: Oxford, Oxford University Press (in press).
- Strabo, The Geography, Translated by H.L. Jones, 1917–1932: Cambridge, Harvard University Press, 8 volumes.
- Todt, W., Cliff, R.A., Hanser, A., and Hofmann, A.W., 1996, Evaluation of a ²⁰²Pb-²⁰⁵Pb double spike for high-precision lead isotope analysis, in Hart, S.R., and Basu, A., eds., Earth Processes: Reading the Isotope Code: Washington D.C., American Geophysical Union, p. 429–437.
- Tufnell, O., 1953, Lachish III: The Iron Age, text and plates: London, Oxford University Press, 437 p.
- Véron, A., Goiran, J.P., Morhange, C., Marriner, N., and Empereur, J.Y., 2006, Pollutant lead reveals the pre-Hellenistic occupation and ancient growth of Alexandria, Egypt: Geophysical Research Letters, v. 33, L06409, doi: 10.1029/2006GL025824.
- Wali, A.W.A., Brookfield, M.E., and Schreiber, B.C., 1994, The depositional and diagenetic evolution of the coastal ridges of northwestern Egypt: Sedimentary Geology, v. 90, p. 113–136, doi: 10.1016/0037-0738(94)90020-5.
- Warne, A.G., and Stanley, J.-D., 1993, Late Quaternary evolution of the northwestern Nile delta and adjacent coast in the Alexandria region, Egypt: Journal of Coastal Research, v. 9, p. 26–64.
- Weill, R., 1919, Les ports antéhelléniques de la côte d'Alexandrie et l'empire crétois: Bulletin de l'Institut Français d'Archéologie Orientale, v. 16, p. 1–37.

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